STABILITY ANALYSIS OF UNFAVOURABLE GEOLOGICAL BODIES IN A POST-EARTHQUAKE AREA: A CASE STUDY IN GENGDA TOWNSHIP CHINA

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Abstract. On 12 May 2008, a devastating mega-earthquake of magnitude 8.0 struck the Wenchuan area, northwestern Sichuan Province, China. Following this earthquake event, many areas in the affected areas were susceptible to geological disasters, such as debris flows, landslides and other secondary disasters. To better understand the mechanism of the formation processes of geological disasters, and to reduce the economic loss caused by the disasters, a comprehensive analysis is required. With the aim of obtaining the characteristics of unfavourable geological bodies, and analysing their stability, we choose the Gengda Township as an example (there are many unfavourable geological bodies distributed in this area, including landslides and deformation bodies). Field investigations and computational models are employed to analyse the stability of the unfavourable geological bodies. Through preliminary analysis and model calculation, the following conclusions can be drawn: under natural conditions, the stability factor of the landslide ranges from about 1.02-1.03, which means that the landslide is in a basically stable or less stable condition. Under the effect of continuous heavy rainfall, the stability factor of the landslide will be significantly reduced, and the corresponding value ranges from 0.9–0.96, which signifies that the landslide is in an unstable condition. For No 1 and No 2 deformation bodies (named by local inhabitants), the stability varied from less stable to unstable, and basically stable to unstable, respectively. On the basis of the results of the calculation, some countermeasures are proposed to mitigate the effects of these disasters probably caused by geological bodies.

Keywords: case study, stability, model calculation, safety factor, transfer coefficient method

Introduction

The 2008 Wenchuan Earthquake in Sichuan Province, China generated many unfavourable geological bodies, which can provide a huge amount of loose deposits, which have caused a dramatic increase in debris-flow occurrence in subsequent years (Hu et al., 2016; Tang et al., 2009; Zhuang et al., 2013). The epicenter Wenchuan Earthquake was in Wenchuan County, Sichuan (31°1'15.60"N, 103°22'1.20"E). As a consequence of the severe shocks, most valleys and slopes were destabilized and

numerous geo-hazards, such as landslides, collapses, and unstable slopes, were triggered in the earthquake-affected area, which covered 50 badly affected counties and ten worst-affected counties in Sichuan province, SW China (Chen et al., 2009; Lee et al., 2017; Wang et al., 2017). The earthquake-stricken area is also the key area for torrent and debris flow mitigation in Sichuan. It was observed that the availability of masses of loose materials was the most important contributing factor for debris flows. Therefore, rock avalanches and landslides induced by the earthquake would supply abundant solid materials to form subsequent debris flows (Chen et al., 1989; Jiayuan, 1995; Postance et al., 2017; Sun et al., 2016). Gengda township is located in the worststricken area, Wenchuan earthquake not only caused a lot of casualties and economic losses, but also induced numerous secondary geological disasters, which including collapses, landslides and debris flows (Ni et al., 2012). As triggered by the Wenchuan earthquake, abundant rock falls and landslides were deposited in the gully drainages, which could provide a large amount of loose solid materials that might be directly turned into debris flows or easily entrained by debris flows. The abundant loose solid materials produced by an earthquake is usually recognized as the most important factor contributing to the occurrence of debris flow. Furthermore, the threshold of trigger factors, such as rainfall, significantly decreases after earthquake events. Thus, areas affected by the Wenchuan earthquake were prone to geological disasters (Ni et al., 2011; Zhuang et al., 2010; Kang and Kim, 2016; Shima et al., 2016).

Through field investigations, the lithology, seismic activity, and tectonic movements of the study area can be determined and consist of the following: exposed strata mainly concentrated in the Quaternary alluvial layer (Q_4^{al+pl}) , the eluvial layer (Q_4^{el+dl}) , colluvium (Q_4^{col}) , and the landslide deposition layer (Q_4^{del}) . Moreover, the bedrock in the study area is fragmentized due to the effect of tectonic movement and weathering, which can greatly facilitate landslide movement in subsequent rainfall conditions.

The study area is located in Huaxia tectonic belt, which belongs to Longmenshan tectonic system, and the tectonic movement is extremely complex. Due to the effect of strong tectonic movements, the bedrock becomes loose and can thus provide favourable conditions for the development of collapse, landslides, and other geological disasters.

On 12 May 2008, a devastating mega-earthquake of magnitude 8.0 struck the Wenchuan area of northwestern Sichuan Province, China. The focal mechanism of the earthquake was successive massive rock fracturing at 15 km depth at Yingxiu. Seismic analysis confirmed that the major shock occurred on the Beichuan–Yingxiu Fault and that aftershocks rapidly extended in a straight northeast–southeast direction along the Longmenshan Fault zone. The total number of fatalities approached 15000, with a significant number resulting from seismically triggered geohazards (Cui et al., 2011). The study area is located in Longmenshan seismic zone. Several earthquakes hit this area before the 12 May 2008 Wenchuan earthquake, including the 1933 Diexi earthquake (with a magnitude of 7.5), the 1657 Wenchuan earthquake (with a magnitude of 7.2). The study area influenced by these earthquakes reached degree V.

In order to mitigate the disasters that were caused by unfavourable geological bodies after the Wenchuan earthquake, field investigations and calculation models are employed in this paper to analyse the stability of the unfavourable bodies. After analysis, the stability of the unfavourable bodies can be identified, and the corresponding engineering countermeasures can be proposed. Thus, the magnitude of the disasters caused by the unfavourable bodies can be significantly reduced, and the security of the inhabitants living adjacent to the unfavourable bodies can be guaranteed.

Materials and methods

Study area

Wenchuan earthquake of 12 May 2008 which caused numerous coseismic landslides. After earthquakes of this magnitude, most valleys and slopes are destabilized and conditions for debris flows and landslides are amplified. These hazards are typically very active during the following 10-20 years. As a consequence of the severe shocks, most valleys and slopes were destabilized and numerous geo-hazards, such as landslides (Fig. 1), collapses (Fig. 2a), and unstable slopes (Fig. 2b), were triggered in the earthquake-affected area, which covered 50 badly affected counties and ten worstaffected counties in Sichuan province, SW China (Chen et al., 2009). The Gengda township is located in an earthquake-stricken area; due to the effect of seismic activity, many slopes located in the Gengda township are seriously affected. Thus, in the subsequent rainy seasons, these unfavourable bodies can be transferred into secondary disasters in response to heavy rainfall. The main objective of this paper is twofold. First, through field investigation, the fundamental data of the study area and the distribution of the unfavourable bodies can be identified; then, computational models are employed to compute the safety factor of each slope, which could allow us to analyse the stability of the slopes distributed within the study area. On the basis of the model calculations, the risk of each slope can be mapped, and preventative engineering can be designed. Thus, the threat caused by the unfavourable bodies can be greatly reduced.

Characteristics of the unfavourable geological bodies

The unfavourable geological bodies in the Gengda township mainly fall into the following four types: landslides, slippery bodies, deformation bodies, and perilous rock. These unfavourable geological bodies cause a great risk to the security of the nearby inhabitants. Moreover, many facilities are also under their threat (*Table 1*). In order to compute the safety factors of the unfavourable geological bodies, field investigations were conducted meticulously. Through these investigations, the fundamental data of the study area could be obtained, including meteorological data, geological conditions, topography, rainfall data, tectonic and earthquake history, hydrological conditions, and human activity. With the help of the data obtained through field investigations, the stability of the bodies can be analysed preliminarily. Then, calculation models are employed to compute the corresponding safety factors of he bodies. By comparing the calculation results to preliminary analysis results, the stability of the unfavourable geological bodies can be reasonably analysed.

Methods

In this paper, two basic approaches (field investigations and model calculations) are employed to explore the stability of the unfavourable bodies. The field investigations can help us acquire data that are closely related with the unstable bodies, which will be useful and can greatly facilitate the subsequent model calculation processes. Based on the fundamental data of the study area and the results of model calculations, a comprehensive stability analysis of each body can be performed.

Field investigations

The field investigations were conducted from the period between June 2009 and July 2009, and these work lasted for around one month. Our investigations are primarily concentrate in Gengda township, Wenchuan prefecture, Sichuan Province. During the processes of investigations, we mainly collected the data with regard to rainfall conditions, geological conditions, topography, etc. The main purpose of the field investigations concerned the following main issues: an initial consideration of the distribution, scale, geological conditions, and triggering factors of each unstable body, followed by a preliminary analysis of the formation mechanism of each unstable body alongside an evaluation of its stability. These results can be compared with subsequent model computations, which can help us to draw reasonable and accurate conclusions. These conclusions can provide a reliable geological basis for designing preventative engineering measures for each unstable body. With the help of Wolong Reserve Management Bureau, Wolong Land Resources Bureau, and Wolong township government, field investigations were conducted and lasted for around one month. The scope of the investigations not only concentrated on the unstable body itself but also included the adjacent zone (within a 20 metre radius) and the area susceptible to these unstable bodies. Through the field investigations, it was possible to acquire the fundamental data of the study area, which are described in the following sections.



Figure 1. Full view of Nitianguo landslide (red point is the study area). a Landslide range, b location of the study area

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Figure 2. Other types of unstable geological bodies. a Collapses, b unstable slopes

Table	1.	Number	of	buildings	and	inhabitants	under	the	threat	of	unfavourable	geological
bodies	in	Gengda	tov	vnship								

Object under threat of unfavourable geological bodies	Name of household	Inhabitants	Area of building (m ²)	Potential economic loss (10 ⁴)	Notes
	Shunfang Liu	5	230	10	
	Qing Jian	4	300	20	Original building was damaged
	Xing Jian	4	140	8	
	Lin Jian	5	380	30	
	Fugui Yu	4	400	12	
	Tianguo Ni	5	500	180	
	Shaoxiu Zhou	10	600	20	
	Tianyuan Ni	5	400	15	
Gengda township	Tianhua Ni	5	150	12	
	Wenxiang Tang	6	150	4	
	Bing Jian	5	120	10	Under construction
	Huachuan Gong Highway Construction Group Co., Ltd. 303 Project Department	50	Rented building	200	
	Workers of Huachuan Gong Highway Construction Group Co., Ltd	50	Rented building	5	
Transmission facilities	Transmission tower			60	
Disaster prevention facilities	Retaining wall (2×8×85 m)		85		Under construction
Highway	Highway 303		500	150	
Number	11 households, 58 people transmission tower	e, floating pop rs and 500 m l	ulation 100, 2 nighway	736	

Model calculations

Since most slopes and other deformable bodies distributed within the study area mainly consist of soil, the deformation of the slope is mainly governed by the weak interface with maximum shear stress. Therefore, we can use the Transfer Coefficient Method (GB 50330-2013) to compute the safety factors of the slopes. Following this method, and to facilitate the subsequent deducing processes, the soil is initially divided vertically into several strips (*Fig. 3*) and each strip is taken as a rigid body while taking the interaction forces between the strips into consideration at the same time. The deducing results can be employed to calculate the safety factor of each strip from the slope, and the safety factor of the whole slope can be determined. On the basis of the above description, we assumed that the residual sliding force was parallel to the sliding surface, and then the driving moment and the resisting moment exerted on individual strips can be calculated, as shown by the stress analysis diagram in *Figure 3*.



Figure 3. Calculation model of transfer coefficient method

(1) The computational formula for computing the safety factor is expressed as follows:

$$K_{f} = \frac{\sum_{i=1}^{n-1} ((W_{i}(1-r_{u})\cos\theta_{i})tg\phi_{i} + C_{i}L_{i}\prod_{j=i}^{n-1}\psi_{j}) + R_{n}}{\sum_{i=1}^{n-1} (W_{i}(\sin\theta_{i} + A\cos\theta_{i})\prod_{j=i}^{n-1}\psi_{j}) + T_{n}}$$
(Eq.1)

$$R_i = (W_i((1 - r_U)\cos\theta_i - A\sin\theta_i) - R_{Di})tg\phi_i + C_iL_i$$
(Eq.2)

$$T_i = W_i (\sin \theta_i + A \cos \theta_i) + T_{Di}$$
(Eq.3)

$$\prod_{j=i}^{n-1} \psi_j = \psi_i \psi_{i+1} \cdots \psi_{n-1}$$
 (Eq.4)

APPLIED ECOLOGY AND ENVIRONMENTAL RESEARCH 17(4): 7807-7819. http://www.aloki.hu • ISSN 1589 1623 (Print) • ISSN 1785 0037 (Online) DOI: http://dx.doi.org/10.15666/aeer/1704_78077819 © 2019, ALÖKI Kft., Budapest, Hungary Ψ_i is the residual sliding force transferred from strip i to strip i + 1 (j = i); that is:

$$\Psi_i = \cos(\theta_i - \theta_{i+1}) - \sin(\theta_i - \theta_{i+1}) \tan \phi_{i+1}$$
(Eq.5)

where W_i is the weight of the i-th strip; T_i is the sliding force of the i-th strip; R_i is the resistance force of the i-th strip; θ_i is the inclination of the sliding surface of the i-th strip; β_i is the angle between the sliding surface and the ground water (here, ground water denotes the water level below the i-th strip); Ψ_i is the cohesion forces of the i-th strip; ϕ_i is the internal friction angle of the i-th strip; L_i is the length of the sliding surface of i-th strip; A is the coefficient of earthquake acceleration in the area where the seismic fortification grade is level 6, A is equal to 0.05; R_{D_i} is the component of seepage force perpendicular to the sliding surface, and R_{D_i} is equal to $N_{W_i} \tan \beta_i \sin(\alpha_i - \beta_i)$; T_{D_i} is the component of seepage force parallel to the sliding surface, and T_{D_i} is equal to $N_{W_i} \tan \beta_i \cos(\alpha_i - \beta_i)$; N_{W_i} is pore water pressure, and N_{W_i} is equal to $\gamma_W h_{W} L_i$; r_U is the ratio of pore pressure, and $r_U = \frac{v_w \times \gamma_w}{v_s \times \gamma_s} \approx \frac{s_u}{s_i \times 2}$, where v_w is the volume of the strip below the water level, s_i is the total area of the strip, and γ_w and γ_s represent the unit weight of the water and strip (sliding body), respectively.

(2) The computational formula for computing the residual sliding force is expressed as follows:

$$P_{i} = P_{i-1}\Psi_{i-1} + K_{s} \times T_{i} - R_{i}$$
(Eq.6)

where P_i is the sliding thrust of the i-th strip (kN/m); P_{i-1} is the residual sliding force of i-th strip (kN/m); and K_s is the designed safety factor, with a value of 1.15 here.

Climatic and geological settings

Climatically, the study area falls into the subtropical humid climate zone; during winter the area is significantly affected by the Qinghai-Tibet Plateau climate, while in summer it is mainly dominated by the southwest and southeast monsoon climate. The mean annual precipitation is 930.2 mm, and the rainfall is mainly concentrated during the period from May to September, which contributes more than 78% of the total precipitation throughout the year. According to the rainfall contour diagram of Sichuan province illustrated in the manual of flood calculations for small and medium-sized basins, the rainfall data for 10 min and 1, 6, and 24 h can be acquired and the corresponding amounts are 10, 20, 65, and 120 mm, respectively. Some slopes in the study area have been loosened by the effect of seismic activity, which will result in instability under subsequent heavy rainfall conditions (Tang et al., 2011). The rainfall data from 2000 to 2009 were gathered, as shown in *Table 2*, and the historical rainfall histogram is shown in *Figure 4*.

	Rainfall (mm)												
Year	Jan	Feb	Mar	Apr	May	June	July	Aug	Sep	Oct	Nov	Dec	rainfall (mm)
2000	8.0	13.0	30.0	122.0	144.0	122.0	/	/	29.0	78.0	46.0	1.0	
2001	8.0	11.0	26.0	89.0	94.0	154.0	161.0	165.0	149.0	111.0	16.0	9.0	993
2002	4.0	24.0	46.0	68.0	155.0	244.0	303.0	115.0	114.0	92.0	12.0	0	1177
2003	7.0	11.0	32.0	82.0	96.0	159.0	207.0	139.0	123.0	91.0	15.0	2.0	964
2004	6.0	11.0	24.0	117.0	97.0	189.0	152.0	106.0	168.0	73.0	28.0	6.0	977
2005	/	/	/	24.0	133.0	160.0	235.0	184.0	99.0	34.0	/	/	
2006	/	/	/	11.0	113.0	95.0	152.0	118.0	43.0	36.0	9.0	2.0	
2007	4.0	33.0	36.0	83.0	338.0	146.0	132.0	9.0	/	/	/	/	
2008	7.0	23.0	32.0	81.0	113.0	103.0	218.0	186.0	98.0	118.0	35.0	5.0	1019
2009	4.0	14.0	16.0	132.0	141.0	140.0	195.0	177.0	132.0	32.0	41.0	20.0	1044

Table 2. Rainfall data from 2000 to 2009 (the dataset are collected from county annals)



Determination of the load condition parameters for model calculations

Through field investigations, the influential factors of the unstable bodies in the study area were studied meticulously. By analysis we found that the slopes in this area are mainly affected by the own weight of the slope, rainfall, and seismic activity. Based on these influential factors, we determined the load conditions and computed the safety factors under different load conditions. The load conditions were mainly grouped into three types: condition 1 represents the effect of the own weight of the slope and rainfall conditions (34 mm/h), and condition 3 represents the effect under the own weight of the slope, rainfall conditions (34 mm/h), and seismic activity.

During field investigations, we conducted some in-situ bulk density tests (we dug a hole in the field, measured the size of the hole in order to calculate the volume of the soil, and weighing the weight of the soil that was dug out, thus the unit weight can be figured out. Use similar method, we can also determine the saturated unit weight of the soil), and obtained the corresponding data, then combined with local rock-mass unit weight the unit weight of the slope body can be determined, the value of unit weight and saturated unit weight of the slope body is 17.1 and 25 kN/m³, respectively. Soil strength

parameters can be obtained by laboratory experiments (direct shear tests in the laboratory were conducted to identify strength of the soil collected from study area), by experiments the shear strength under fast shear and under fully saturated shear can be acquired simultaneously, their corresponding value shown in *Table 3*.

Parameters	Shear strengt conc	h under natural litions	Shear strength under fully saturated conditions			
	C (kPa)	Φ (°)	C (kPa)	Φ (°)		
Sliding body and No 1 deformation body	8.5	29	7.5	28		
No 2 deformation body	3.5	23.5	2.5	22		

Table 3. Shear strength of sliding body under different conditions

Results and discussion

According to the principle of the Transfer Coefficient Method, the unfavourable bodies in the study area can be vertically divided into several strips. Through preliminary analysis, we found that strips 6 and 7 are the most dangerous sliding bodies in the whole landslide, strips 1 and 2 are the most dangerous sliding bodies in No 1 deformation body, and strip 5 is the most dangerous body in No 2 deformation body. The stability calculation diagrams are shown in *Figure 5a-c*.



Figure 5. Stability calculation diagram of a strip 1 of No 1 deformation body (profile); b strip 2 of No 1 deformation body (profile) and c strip 5 of No 2 deformation body (profile)

On the basis of Section 4, the calculation can be conducted and the calculation results can be obtained. Then the safety factors of strips 6 and 7 of the landslide, strips 1 and 2 of No 1 deformation body and strip 5 of No 2 deformation body can be obtained under different load conditions, as shown in *Table 4*. For No 1 deformation body, under the influence of the own weight of sliding body, the safety factor is 1.04, and under the effects of the own weight of sliding body and rainfall (34 mm/h), the safety factor is 1.03, results of the analysis can also be applied to landslide and No 2 deformation body.

Table 4. Model calculation results. Conditions 1, 2 and 3 represent the effects under the own weight of landslide, the effects under the own weight of landslide and rainfall (34 mm/h), the effects under the own weight of landslide, rainfall (34 mm/h), and seismic activity

I and conditions unstal	le hodies and string	1	2	3	
Loau conunions unstat	ble boules and strips	Safety factor	Safety factor	Safety factor	
Landelida	6	1.02	0.9	0.61	
Lanushue	7	1.13	0.96	0.66	
No.1. defermention hadre	1	1.04	0.89	0.62	
No 1 deformation body	2	1.03	0.89	0.63	
No 2 deformation body	5	1.1	0.92	0.60	

According to "Landslide Prevention Engineering Design and Construction of Calculation Specification", on the basis of the classification of landslide stability, the stability state under different stability factors can be determined, as shown in *Table 5*. With the help of *Table 5*, the safety factors under three different load conditions can be set as 1.30, 1.10, and 1.05, respectively. Then the residual sliding thrust can be calculated, and the calculation results are shown in *Table 6*.

Table 5. Classification of landslide stability (Kst is the designed safety factor) Conditions 1, 2 and 3 represent the effects under the own weight of landslide, the effects under the own weight of landslide and rainfall (34 mm/h), the effects under the own weight of landslide, rainfall (34 mm/h), and seismic activity

Stability factor of the slope	$K_{s} < 1.00$	$1.00 < K_s \le 1.05$	$1.05 < K_s \leq K_{st}$	$K_s \!\geq\! K_{st}$	
Stability status	Unstable	Less stable	Basically stable	Stable	

According to landslide engineering investigation specifications, the stability of landslides can be subdivided into four classes. If the stability factor of a landslide under certain load conditions is larger than the safety factor, we can state that the stability of the landslide matches the stability criteria well. Based on this principle, the following conclusions can be drawn:

Under natural conditions, the stability factor of strip 6 is 1.02; thus referring to *Table 5*, the landslide is in the less stable state. Through field investigations, we found that many cracks developed at the location of strip 6. Moreover, under continuous rainfall, the stability factor decreased to 0.9, which means that the landslide is unstable under such conditions, which is in good agreement with the preliminary analysis. Under natural conditions, the stability factor of strip 6 is 1.13; referring to *Table 5*, the

landslide is in the basically stable state. Similarly, under continuous rainfall, the stability factor decreased to 0.96, i.e. the landslide is unstable under such conditions, which is in good agreement with the real situation.

Unfavourable be and strips	odies	Load conditions	Safety factor F_s Stability factor F_{st}		Residual sliding thrust E(kN/m)
		Condition 1	1.02	1.2	479.22
	No.6	Condition 2	0.9	1.10	800.28
Londalida		Condition 3	0.61	1.05	2230.20
Landshue		Condition 1	1.13	1.2	87.77
	No.7	Condition 2	0.96	1.10	274.38
		Condition 3	0.66	1.05	945.36
		Condition 1	1.04	1.2	222.72
	1-1'	Condition 2	0.89	1.10	433.78
No 1 deformation		Condition 3	0.62	1.05	1139.49
body		Condition 1	1.03	1.2	348.58
	2-2'	Condition 2	0.89	1.10	634.06
		Condition 3	0.63	1.05	1580.44
		Condition 1	1.10	1.2	55.34
No 2 deformation	5-5'	Condition 2	0.92	1.10	147.33
		Condition 3	0.60	1.05	511.87

Table 6. Calculation results of the residual sliding thrust

Under natural conditions, the safety factors of strips 1 and 2 of No 1 deformation body are 1.04 and 1.03, respectively. Referring to *Table 5*, the No 1 deformation body is in the less stable state. These findings agreed well with preliminary analysis. Under continuous rainfall, the stability factors of strips 1 and 2 both decreased to 0.89, meaning that the No 1 deformation body is unstable under such conditions, which is in good agreement with the real situation.

Under natural conditions, the safety factor of strip 5 of No 2 deformation body is 1.10. Referring to *Table 5*, the No 2 deformation body is in the basically stable state. This agreed well with the preliminary analysis. Under continuous rainfall, the stability factor of strip 5 decreased to 0.92, meaning that the No 2 deformation body is unstable under such conditions, which is in good agreement with the real situation.

Based on this analysis, the following conclusions can be drawn. The stability of the landslide and deformation bodies varied as the load conditions changed. From condition 1 (i.e. under the effect of the own weight of the slope) to condition 3 (i.e. under the effect of the own weight of the slope, rainfall, and seismic activity), the stability of the landslide varied from basically stable to less stable and finally to unstable; the stability of the No 1 deformation body varied from less stable to unstable; and the stability of the No 2 deformation body varied from basically stable to unstable. Due to the discrepancy of the slope in the study area, the deformation and stability variations have some differences; under some circumstances, partial failure may occur in the stable state.

Conclusions

By means of field investigation ands model calculations, the following conclusions can be drawn. The stability of the unfavourable bodies was mainly affected by the 12 May 2008 Wenchuan Earthquake and subsequent rainfall infiltration. Continuous heavy rainfall plays a crucial role in slope stability analysis. In addition, human activity and weathering have also influenced the stability of unfavourable geological bodies in the study area.

Under natural conditions, the stability factor of the landslide ranges from about 1.02–1.03, which means that this landslide is in the basically stable or less stable condition. Under the effect of continuous heavy rainfall, the stability factor of the landslide will be significantly reduced, and the corresponding value ranges from 0.9–0.96, which means that the landslide is in the unstable condition. For No 1 and No 2 deformation bodies, the stability varied from less stable to unstable, and basically stable to unstable, respectively.

At present, some of the unfavourable bodies are in the less stable state, under the effect of continuous heavy rainfall, seismic activity, or other influential factors; some geological bodies will probably change to the unstable state. Considering the property of the landslide itself, a cut slope is recommended as a preventative and mitigation measure. For the No 1 deformation body, anti-slide piles are recommended as prevention and mitigation measures. Due to the small volume and simplicity of the No 2 deformation body, it is recommended to clear it directly.

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REFERENCES

- [1] Chen, B., Wei, L., Wu, C. (1989): Analysis on the method of pattern recognition on the development, distribution and the restrictive elements of debris flows in southwestern China. Bulletin of Soil and Water Conservation 9(2): 54-56.
- [2] Chen, N., Yang, C., Zhou, W., Hu, G., Li, H., Hand, D. (2009): The critical rainfall characteristics for torrents and debris flows in the Wenchuan earthquake stricken area. Journal of Mountain Science 6(4): 362-372.
- [3] Chuan, T., Jing, Z., Xin, Q., Jun, D. (2011): Landslides induced by the Wenchuan earthquake and the subsequent strong rainfall event. A case study in the Beichuan area of China. Engineering Geology 122: 22-33.
- [4] Cui, P., Chen, X.-Q., Zhu, Y.-Y., Su, F.-H., Wei, F.-Q., Han, Y.-S., Liu, H.-J., Zhuang, J.-Q. (2011): The Wenchuan earthquake (May 12, 2008), Sichuan province, China, and resulting geohazards. – Natural Hazards 56(1): 19-36.
- [5] GB 50330-2013 (2013): Technical Code for Building Slope Engineering. National Standard of People's Republic of China.
- [6] Hu, W., Dong, X. J., Xu, Q., Wang, G. H., van Asch, T. W. J., Hicher, P. Y. (2016): Initiation processes for run-off generated debris flows in the Wenchuan earthquake area of China. – Geomorphology 253: 468-477.
- [7] Jiayuan, T. A. O. (1995): The distribution of disastrous debris flow in China. Journal of Higher Correspondence Education (Natural Science) 4: 6-7.
- [8] Kang, H.-S., Kim, Y.-T. (2016): The physical vulnerability of different types of building structure to debris flow events. Natural Hazards 80(3): 1475-1493.

- [9] Lee, C. F., Huang, W. K., Chang, Y. L., Chi, S. Y., Liao, W. C. (2017): Regional landslide susceptibility assessment using multi-stage remote sensing data along the coastal range highway in northeastern Taiwan. Geomorphology 300: 113-127.
- [10] Ni, H. Y., Zheng, W. M., Tie, Y. B., Su, P. C., Tang, Y. Q., Xu, R. G., Wang, D. W., Chen, X. Y. (2011): Formation and characteristics of post-earthquake debris flow: a case study from Wenjia gully in Mianzhu, Sichuan, SW China. – Natural Hazards 61(2): 317-335.
- [11] Ni, H., Zheng, W., Tie, Y., Su, P., Tang, Y., Xu, R., Wang, D., Chen, X. (2012): Formation and characteristics of post-earthquake debris flow: a case study from Wenjia gully in Mianzhu, Sichuan, SW China. – Natural Hazards 61(2): 317-335.
- [12] Postance, B., Hillier, J., Dijkstra, T., Dixon, N. (2017): Comparing threshold definition techniques for rainfall-induced landslides: a national assessment using radar rainfall. – Earth Surface Processes & Landforms. DOI: 10.1002/esp.4202.
- [13] Shima, J., Moriyama, H., Kokuryo, H., Ishikawa, N., Mizuyama, T. (2016): Prevention and mitigation of debris flow hazards by using steel open-type Sabo dams. – International Journal of Erosion Control Engineering 9(3): 135-144.
- [14] Sun, G., Huang, Y., Li, C., Zheng, H. (2016): Formation mechanism, deformation characteristics and stability analysis of Wujiang landslide near Centianhe reservoir dam. – Engineering Geology 211: 27-38.
- [15] Tang, C., Zhu, J., Li, W., Liang, J. (2009): Rainfall-triggered debris flows following the Wenchuan earthquake. – Bulletin of Engineering Geology and the Environment 68(2): 187-194.
- [16] Wang, J., Yang, S., Ou, G., Gong, Q., Yuan, S. (2017): Debris flow hazard assessment by combining numerical simulation and land utilization. – Bulletin of Engineering Geology and the Environment77(1): 13-27.
- [17] Wang, Y., Zhao, B., Li, J. (2017): Mechanism of the catastrophic June 2017 landslide at Xinmo Village, Songping River, Sichuan Province, China. – Landslides 4: 1-13.
- [18] Zhuang, J., Cui, P., Hu, K., Chen, X., Ge, Y. (2010): Characteristics of earthquaketriggered landslides and post-earthquake debris flows in Beichuan County. – Journal of Mountain Science 7(3): 246-254.
- [19] Zhuang, J.-Q., Cui, P., Peng, J.-B., Hu, K.-H., Iqbal, J. (2013): Initiation process of debris flows on different slopes due to surface flow and trigger-specific strategies for mitigating post-earthquake in old Beichuan County, China. – Environmental Earth Sciences 68(5): 1391-1403.