

THE HYDROLOGICAL INFLUENCE OF FOREST HARVESTING INTENSITY ON STREAMS: A GLOBAL SYNTHESIS WITH IMPLICATIONS FOR POLICY

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Abstract. The hydrological properties of the clearcutting of forested catchments were widely investigated by analyzing runoff in the pre- and post-harvesting periods. Deforestation worldwide is primarily to meet the wood and fiber products demand for household and industry. It is a widely known phenomenon that deforestation enhances the streamflow and water yield. However, due to the complexity of forest structure and functions, little is known about the exact estimation of a percent increase in water yield after various harvesting intensities of conifers and broadleaved forest globally. To assess these effects, this study analyzed 145 catchments dataset collected from 21 publications. The study evaluates the influence of 25, 50, 75, and 100% deforestation on streamflow. Moreover, changes in the context of various variables like treatment years, elevation, area and mean annual precipitation were also analyzed. Overall comparison showed that after harvesting of broadleaved water yield increases up to 8-23% and in needle-leaved up to 9-28%. The study provides scientific insight into the essential role that annual precipitation, area, elevation, and year of treatment play in influencing hydrology. This research suggests that a target specific approach should be adopted in future forest management under the umbrella of integrated research to mitigate the challenges of climate change.

Keywords: *climate change, streamflow, annual precipitation, water yield, broadleaved, needle-leaved*

Introduction

The ever-increasing trend in the human population has caused an upsurge the overexploitation of natural resources, especially the degradation of forests in terrestrial ecosystems. Forests are essential to life on Earth, providing numerous ecosystem services (Costanza et al., 1997) such as fruits, honey, oil, pickle, biocontrol, pollination, Carbon sink, water, and nutrient recycling as well as biodiversity conservation (Nasi et al., 2002; Badshah et al., 2017; Wang et al., 2017; Masiero et al., 2019; Ullah et al., 2019b; Muhammad et al., 2020). Among all the forest ecosystem services, carbon sinks and water provision to down- stream are the two primary services which act as essential cogs in the carbon and water cycle by playing their active role in forest processes and functions. However, there are some trade-offs between gain in forest productivity and ecosystem water balance (Farooqi et al., 2019b). Due to brimming of the population in the world has led to an increase in anthropogenic disturbance which are the primary cause of changes in forest composition over a period characterized by drastic changes in both land use and cover resulting increase in fossil fuel emissions and influencing environmental condition (Law et al., 2002; Houghton, 2012; Siddique et al., 2020). This situation is getting worst in

developing countries because of massive deforestation and fire incidences (Khan et al., 2019; Ullah et al., 2019a; Ali et al., 2020).

Despite all the conflicting debates on retaining and removing forests (Popkin, 2019), for sequestering carbon which enhances productivity, it is overwhelmingly considered the top priority (Krankina et al., 1997; Ruddell et al., 2007). As far as their interaction with water is concerned in this modern era, forests are also recognized in two important terminologies “upstream” as a source of water in streams and rivers (Zhang et al., 2017), and “upwind” as a source of precipitation (van der Ent et al., 2010; Ellison et al., 2012, 2017) however, these trends are bound to the localities and regions.

Deforestation is mainly considered as a positive aspect of increasing the streamflow and runoff, which is ultimately utilized by the industry and household (Bosch and Hewlett, 1982; Jones and Post, 2004). Meanwhile, the expansion of forests reduces this water flow leading to 52% of half dryness and 13% of complete dryness of streams in the world (Andréassian, 2004; Jackson et al., 2005). The result is decline in water availability to downstream users (van Dijk and Keenan, 2007), especially dry areas are more vulnerable to this situation. However, the phenomenon of annual runoff is generally dependent on annual precipitation and evapotranspiration. The greater the precipitation, the less evapotranspiration will ultimately enhance runoff and vice versa (Komatsu et al., 2011). The proportionate contribution of precipitation to streamflow varies by how interception and evapotranspiration are influenced by vegetation development stage, rooting depth and health. However, this may differ widely according to vegetation type (Calder, 1999; Zhang et al., 2001). Because the main components of evapotranspiration are canopy transpiration and interception loss (Van Wijk et al., 2001; Vertessy et al., 2001; Wilson et al., 2001). Interception losses from coniferous and broadleaved forests were presented by (Huber and Iroume, 2001; Komatsu et al., 2011), depending on rainfall and forest characteristics (Iroume and Huber, 2002).

In the past, many research investigations have evaluated the effect of logging operation on the global variation in water yield depending on different forest types and structure (Hornbeck et al., 1993; Troendle et al., 2001; Andréassian, 2004; Adams and Flower, 2006; Komatsu et al., 2011), especially the impacts of forest harvesting of broadleaves and conifers forests on runoff and water yield (Komatsu et al., 2011). Yet questions and misconceptions linger regarding the influence of forest harvesting operations on streamflow under the variety of climatic, physiographic factors, and forest management constraints. It has been shown that considerable change in streamflow after forest cutting can be observed when more than 20% of the forest cover declined (Stednick, 1996). However, many previous studies of broadleaf and conifers forests reported that the annual runoff improved by 10-70%, depending on the size of the harvesting intensity (Keppeler and Ziemer, 1990; Fahey, 1994; Swank et al., 2001; Farooqi et al., 2020a). Similarly, some reported the effects of timber removal only in the first years after final harvest (David et al., 1994; Bari et al., 1996), whereas, others investigated up to 6-23 years after the event (Ruprecht and Stoneman, 1993; David, 1994; Fahey, 1994). This variation in results widely depending on forest type, harvesting technique, climate as well as the topography of the area.

Although forest harvesting has positive impacts on streamflow and water yield, it also has many adverse implications on the whole ecosystem. Therefore, predefined knowledge about forest types can be helpful for understanding and implementing afforestation/deforestation programs in the context of minting the balance between forest carbon sequestration and water conservation. This will provide future assistance to regional forestry planning and forest management. In this case, the negative influences of forest on

streamflow might be to control the proportion of forest cover at the catchment scale, which has the potential to modify the streamflow regime (Zhang et al., 2012). This fact is essential to get a better understanding of the affiliation between runoff concerning forest cover proportion (Brown et al., 2013).

To satisfy the rapidly increasing burdens on water supply and other ecosystem services, a practical approach for managing forests (afforestation/deforestation) is needed to achieve the multifunctional benefits. That mainly addresses the tradeoff between carbon sink and water yielding, which is urgently required (*Fig. 1*). The present article aimed to investigate the effects of various degrees of deforestation on the hydrological properties of different streams with forested catchments, as well as the influence of precipitation afterward.

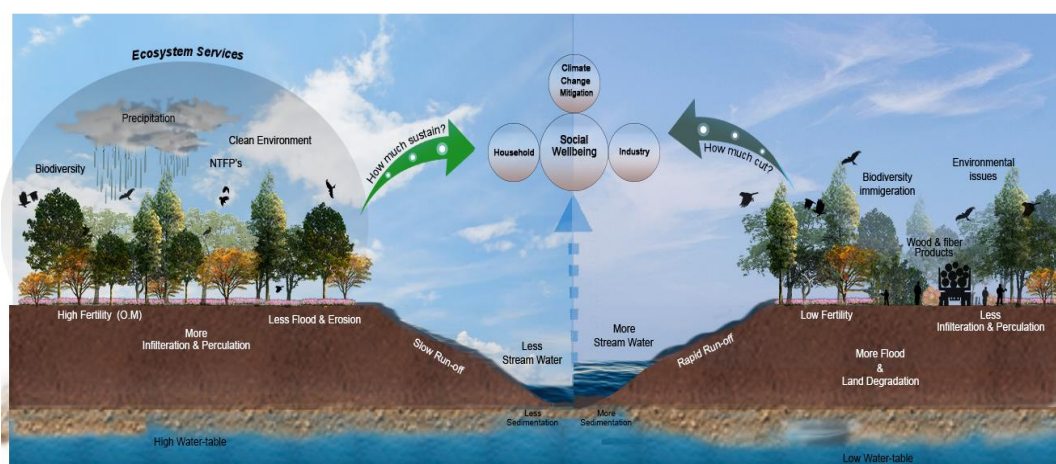


Figure 1. The diagram showing the importance of retaining and removing forest, and their overall impacts on socio-economic development under the umbrella of efficient forest management strategies for climate change mitigation

Materials and methods

Data collection and processing

We have compiled this large dataset of deforestation studies and their impacts on water yield from research articles published peer review journals. The sample consists of total of 64 watershed sites of conifers forest stand and 81 sites belonging to broadleaf forest stand, totaling 300 observations from all over the world. This study compiled the dataset from 21 peer-reviewed journals as well as reports of governmental and nongovernmental research institutes, representing many parts of the world (*Appendix 5*). The forest types were classified into conifers and broadleaf depending upon the dominant species of the forest stand as well as information available in the publication. Information gathering included deforestation intensities on water yield and streamflow before and after treatment. Elevation, age, area, yearly record after treatment and mean annual precipitation were determining from the publication for each site. All those sites which showed no significant increase in water yield after harvesting were discarded to get reliable and expressible estimate of computed harvesting intensities of 25, 50, 75 and 100%. The harvesting intensities were set according to the previous researches guidelines i.e. considerable change in water yield after harvesting was mainly observed when 20% or more area was cut (Bosch and Hewlett, 1982; Stednick, 1996). The percent change in water yield after harvesting was computed with the help of formula as shown in *Equation 1*.

$$\text{Change in water yield (\%)} = \frac{\text{Increase in water yield after treatment}}{\text{Total available water in stream before and after treatment}} * 100 \text{ (Eq.1)}$$

*where stream water before and after treatment in equation.1 is in mm

Testing of significance

First the Normality test i.e. Shapiro-wilk test was performed, this test showed that the conditions of normality and homogeneity of variance were not met and that has been visual represented in QQ plot. Later nonparametric Kruskal–Wallis tests were applied before by Farley et al. (2005) in a kind of synthesis analysis. In each case, the dependent variable was either the proportional change in water yield following change in factors of evaluation i.e. deforestation percentage. The significance test suggests that the water yield rate is not the same in each of the two or more harvesting intensities ($P < 0.05$). Even if we rejected the null hypothesis of no difference, the test does not tell us either the two similar intensities of broadleaved and needle-leaved differ significantly from each other. To compare two groups at a time used the Wilcoxon Rank test.

Results

The results of Shapiro-wilk test rejected normality at $P < 0.0001$ (*Appendix 1*); the results of the QQ plot showed the visual representation, i.e., the distribution of variables for conifers and broadleaves forest groups of all four harvesting intensities. Many points in both ends fall out of the line and are away from the confidence envelope (*Appendix 2*). Similarly, the Kruskal-Wallis test showed a highly significant increase in streamflow (%) after deforestation in broadleaved and conifers forests of the global dataset at $P < 0.0001$ as mentioned in *Appendix 3*. These results reveal that the percent increase in water yield after treatment of 25, 50, 75, and 100% harvesting intensities in the broadleaved forest was 8, 15, 20, and 23%, respectively. However, this increase was significantly higher in conifers than broadleaved with increase of 9, 17, 23 and 28% in water yield respectively (*Figure 2*). Therefore, the overall results of needle-leaved are significantly higher than broadleaved forest stand after treatment as illustrated in *Appendix 3*. Similar results have shown from Wilcoxon rank test while comparing similar harvesting intensities of both the forest vegetation types at ($P < 0.05$) in (*Appendix 4*).

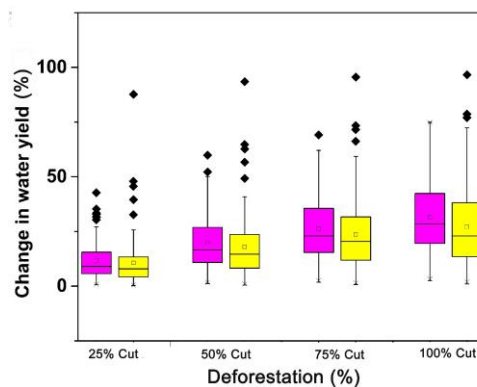


Figure 2. Results showing the percent increase in water yield after different harvesting intensities in broadleaved and needle leaved forests of the world at $P < 0.05$. (Yellow color indicating broadleaved and pink representing needle-leaved forest)

The relationship between annual precipitation and change in water yield (mm) as well as an increase in water yield (%) after harvesting is shown in *Figure 3*. The results showed that in higher annual precipitation regions (>1000 mm), streamflow in mm also increased more than in lower annual precipitation regions (<1000 mm). However, the post-harvest increase in percent change of water yield was higher in the low rainfall area than in high precipitation regions.

The regression analysis in *Figure 4a* and *b* also demonstrated this relationship. The figure illustrated that as long as the annual precipitation (mm) is increasing, the water yield or streamflow (mm) after the treatment also increasing with positive linear trend of $R^2 = 0.35$ at $P < 0.0001$ (*Fig. 4a*). Similarly, the relationship between annual precipitation (mm) and percent increase in water yield or increase in streamflow (%) after treatment showed declining trend with $R^2 = 0.10$, $P < 0.0001$ (*Fig. 4b*), indicating that the percent increase in water yield after treatment was observed from low precipitation to high precipitation regions.

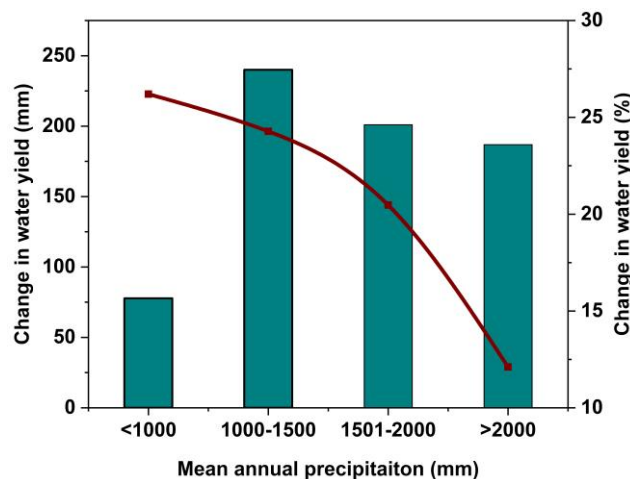


Figure 3. Representing the influences of annual precipitation on change in water yield (mm) and percent change in water yield (%) after harvesting of the study sites (Green bars are representing Change in water yield (mm), brown line showing Change in water yield (%))

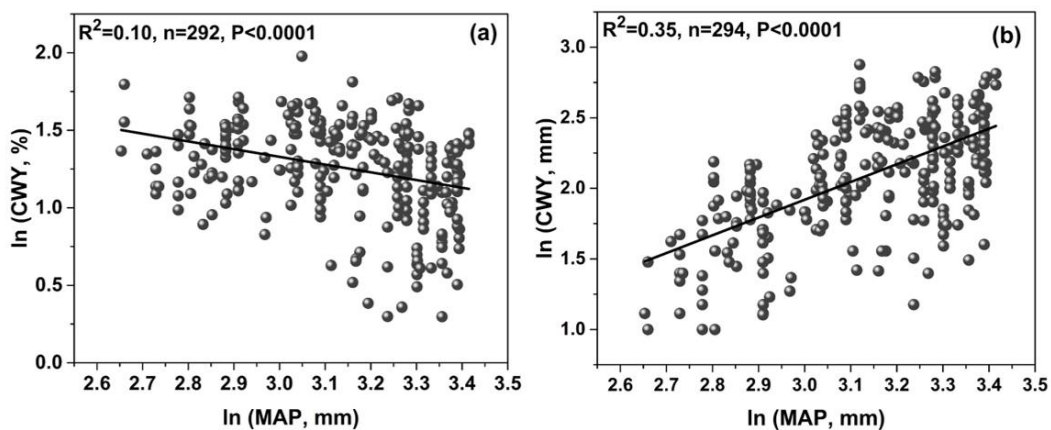


Figure 4. Illustrating the log transformed linear relationship between mean annual precipitation (mm) and change in water yield CWY (mm) and percent change in water yield CWY (%) after treatment of global catchment sites. (MAP-mean annual precipitation)

To further explain, the role of forest types and their interaction with annual precipitation and change in water yield after treatment was assessed in *Figure 5*. The figure indicated that the majority of the broadleaved forests of this study belong to high precipitation areas than needle-leaved forests.

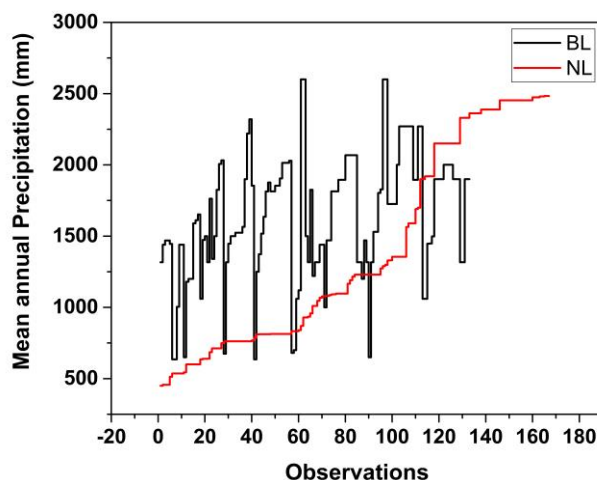


Figure 5. Representing the presence of needle-leaved (NL) and broadleaved (BL) forests observations taken from different precipitation regions of the world in dataset of our study

Similarly, in the dataset majority of the bigger catchment (<150 ha) with higher elevation (<2500 m) were found in lower precipitation regions (>1000 mm) as shown in *Figure 6*.

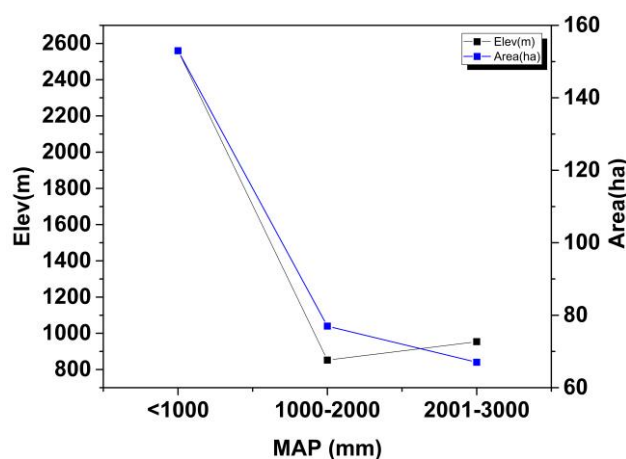


Figure 6. Showing the distribution of mean annual precipitation (mm) at different elevation and forest cover areas of study sites. (MAP-mean annual precipitation)

As far as post-treatment regrowth and recovery of vegetation are concerned, the broadleaf showed significant declining trend at $P < 0.0001$, however these results are non-significant in case of needle-leaved the forests as represented in *Figure 7a* and *b*.

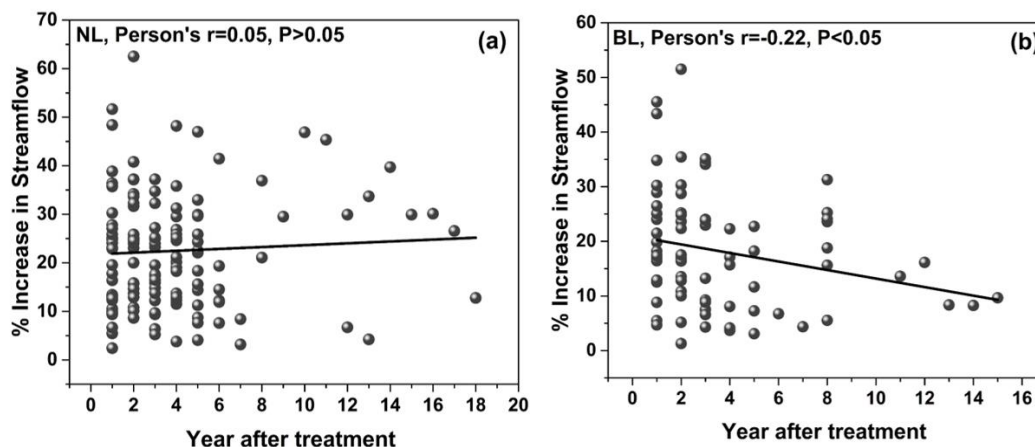


Figure 7. Illustrating the relationship between year after treatment and increase in streamflow (%) of global catchment sites. (NL-needle-leaved, BL-broadleaved forest)

The change in water yield/streamflow after harvesting in mm decrease from lower to higher elevation level (<1000 to 3000 m). However, the percent increase in streamflow after harvesting showed an increasing trend from a lower elevation to higher (<1000 to 3000 m), as shown in *Figure 8*.

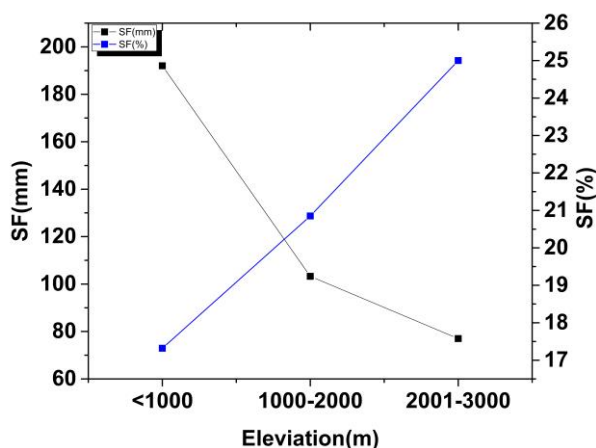


Figure 8. Showing the trend of increase in streamflow (mm) and percent increase in streamflow (%) after harvesting at different level of elevation. (SF-streamflow)

Similarly, in the forest area less than or equal to 100 ha showing 20% increase in streamflow after harvesting, but this trend was at its peak in forest cover of 101-300 ha with maximum percent increase in SF of around 26%, afterword > 300 ha indicating abrupt decline in percent increase of SF up to (18.2%). This is also worth noted that change in water yield in mm is greater (220 mm) in the forests consisting of < 100 ha area followed by decline up to 65-100 mm in forest cover of 101-300 ha land as shown in *Figure 9*.

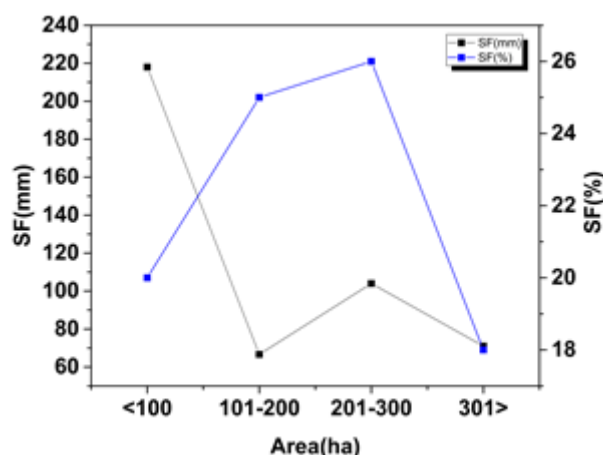


Figure 9. Showing the trend of increase in streamflow (mm) and percent increase in streamflow (%) after harvesting at different forest cover area (ha). (SF-streamflow)

Discussion

Forest types and hydrology

On a global level, there is a significant research gap about exact identification of the increase in water yield (%) and streamflow (mm) of different forest harvesting intensities. However, mixed results of varying harvesting intensities have found in previous research investigations (Bosch and Hewlett, 1982; Hornbeck et al., 1993; Stednick, 1996; Troendle et al., 2001; Pike and Rob, 2003; Andreassian, 2003; Adams and Flower, 2006; Komatsu et al., 2011). The plausible reason behind this variability might be due to different site/location, climate type, vegetation type, forest structure, origin, stand age, treatment years, harvesting technique, season of treatment, soil as well as other methodological and technical constraints. The results of this analysis indicated that needle-leaved forest has resulted a greater change in water yield (%) after harvesting than broadleaved when compared to different forest harvesting intensities of 25-100%. This increase in water yield of broadleaved was (8-23%) and needle-leaved (9-27%) after treatment is shown in *Figure 2*. In the previous research investigations, it was indicated that the considerable change in streamflow after timber harvesting occurred when more than 20% of the forest cover was reduced (Bosch and Hewlett, 1982; Stednick, 1996). However, phenomenon has contradicted and reveals that in some of the catchment studies, lesser harvesting intensity has had measurable increases in water yield than the area with 100% harvest depending on the catchment site and topographic factors. For example, with 15% of the basal catchment area could be cut for a considerable upsurge in annual water yield at the catchment scale in the Rocky Mountain region, whereas 50% in the Central Plains, although system responses are variable (Stednick, 1996). Similarly, the results from previous studies are also in accordance with the findings of this study indicating that the influence of different harvesting intensities on percent change in water yield is higher in needle-leaved than broadleaved forest. A recent global synthesis indicated that 68% removal of broadleaved forest leads to increase of just 16% of stream flow (Farooqi et al., 2020a). Another regional study in New Zealand showed that native deciduous forest clear-felling caused average increase of 70% on five years of treatment (Fahey, 1994) on the

other side in a southern Appalachian Mountains (USA) 59 ha of mixed hardwood stand clearcutting enhanced streamflow just 28% after the first year of treatment (Swank et al., 2001). In two catchment studies in Australia one was patch-cut to remove 22% of basal area of Wicksend catchment, and the Willbob catchment was thinned to remove 12% of basal area of eucalyptus forests. This caused an annual increase in streamflow by 10% in the first three years after logging at Wicksend, and by 31% for the first four years at Willbob (Lane and Mackay, 2001).

Consequently, mixed results have recorded in case of needle-leaved deforestation. For example, the removal of 14 million board feet of lodgepole pine (*Pinus contorta*) from about 25 percent of the Brownie Creek basin formed an average of 147 mm extra water yield per annum, which is equal to 52% of the increase in annual water yield (Burton, 1997). The study of continental/maritime hydroclimatic regions of the United States in naturally regenerated conifers stands after 50% clear cut and 50% partial cut treatments reported increased water yields of 270 mm (36%) and 140 mm (23%) respectively (Hubbart et al., 2007). Similarly, a global study revealed that with 71% deforestation of needle leaved forests caused an increase of 27% in water yield in down streams (Farooqi et al., 2020a). These results agreed that in needle-leaved forest of large coverage >2000ha might produce significant or drastic increase in water yield and increase the risk of severe flooding (Burton, 1997).

Forest types and precipitation

Annual precipitation impacts the scale of water yield intensifications that follow timber harvest operations in forested watersheds (Keppeler and Ziemer, 1990; Brown et al., 2005; Adams and Fowler, 2006; Komatsu et al., 2011). The results in this study indicated the significant increase in annual streamflow (mm) in higher precipitation areas compared to lower regions of the world as shown in *Figure 2* and the linear trend in the relation between annual precipitation and increase in streamflow (mm) is recorded in *Figure 3b*. This shows that the plantation schemes can be successful established in high precipitation region in order to achieve carbon objectives because abundance of water in these regions will not only helpful in enhancing the growth and productivity but also atmospheric circulation. Similarly, the percent increase in water yield after harvesting is lower than in low precipitation regions of the world (*Figs. 2 and 3a*), because the water available in the region is already in sufficient quantity, therefore after harvesting big change even show little difference. Moreover, the more evaporative losses can act positive in enhancing precipitation having sufficient energy to lift the additional atmospheric moisture high enough to condense and form clouds (Jackson et al., 2005). However, the precise estimation of hydrological implications of large watersheds (> 1000 km²) are largely lacking due to more complex for structure and other confounding factors.

It is also worth noting that in this dataset majority of vegetation at comparatively lower precipitation regions is needle-leaved compare to broadleaved found in higher precipitation regions (*Fig. 5*). A recent past, Farooqi et al. (2020a) highlighted the influences of precipitation on percent increase in water yield after-harvesting in broadleaved and conifers forests, however, he did not elaborate on these impacts and their causing factors. The reason behind all of these results might be vegetation affects the proportion of precipitation that is evaporated and transpired and, consequently, the amount available for soil moisture storage, groundwater recharge, and dry weather streamflow of broadleaved and needle-leaved forests (Komatsu et al., 2011). The

variation in transpiration in the forests is because of the leaf area index as well as stomatal conductance (Kelliher et al., 1995; Raupach, 1995), whereas the interception losses also vary according to leaf area index (Komatsu et al., 2008; Muzylo et al., 2009). These interception losses were thoroughly discussed in the previous research investigations of coniferous and broadleaved forests (Huber and Iroume, 2001; Komatsu et al., 2011), while Iroume and Huber (2002) demonstrated that there are many factors associated with these losses influenced by rainfall and forest characteristics like species, density, age, etc. It is generally believed that the streamflow response depends on the mean annual precipitation of the area (Bosch and Hewlett, 1982; Ruprecht and Stoneman, 1993; Iroumé et al., 2000). Increases in streamflow (mm) are generally most significant in areas of high rainfall, but they are short-lived due to rapid regrowth of vegetation (Bosch and Hewlett, 1982; Ruprecht and Stoneman, 1993; David, 1994; Fahey, 1994; Swank et al., 2001). The decreasing trend toward pre-disturbance levels is of interest because regeneration has been reported in diverse environments, silvicultural and forest species dominance (Fahey, 1994; Bosch and Hewlett, 1982; Cornish, 1993; Hornbeck et al., 1993). For example, reductions in streamflow below pre-disturbance levels have been observed as isolated cases in needle-leaved evergreen planted a forest of the temperate region in southern Chile. Indicated that the 120% increase in runoff might be partly due to the higher rainfall during the post-harvesting period (Iroumé et al., 2006). Another study on the jarrah forest in south-western Australia reveals that the subsequent recovery of vegetation cover has led to water yields returning to pre-disturbance levels after an estimated 12-15 years (Ruprecht and Stoneman, 1993). The deciduous conversion to pine, forest harvesting in moderate-to-high rainfall areas causes a 60-80% increase in water yield for three-five years after clear-felling. It was also noted that the yields should return to pre-harvesting levels within six-eight years, depending on the silvicultural regime adopted (Fahey, 1994).

In the present study results, the significant decline trend in water yield ($P < 0.0001$) after the first year of broadleaved forests removal till it reaches to the pretreatment stage as shown in *Figure 6* might be connected to their coppiced nature which might be the reason for rapid regeneration after deforestation. For example, in a study conducted in Central Portugal, when a coppicing a fast-growing species of eucalyptus due to the fast regrowth of the forest stands recorded that the hydrological effects of clearcutting were short-lived (David, 1994), in Coweeta, a mature hardwood coppice stand the first cutting required 23 years' recovery time to reach pretreatment level in striking contrast the second cutting achieved this level just within 16 years (Swank et al., 1970). Therefore, water use strategies were developed according to the variation in developmental stages as well as the available water resources (Su et al., 2014). The result of these studies demonstrate that annual water yield increases obtained from complete forest cutting in coppice catchment can be more short-lived in second rotation. Because of the difference in basal area, LAI, species density as well as litter fall production of first stand enhances fertility and water retention in the soil, which boost the regrowth of second cutting. Moreover, the only way forward of gaining large increase in annual water yield is to manage regrowth and control dense sprouting and rapid crown development.

Forest types and water use

The main distinction between the percent increase in water yield after harvesting in conifers versus broadleaved as shown in *Figure 2* might be due to the efficient water use of broad-leaved than in conifers. Evergreen conifers tend to have a higher water use

due to high interception losses which are maintained throughout the whole year, and particularly during the winter period when conditions are usually wettest and windiest. During the vegetative period, interception rates are also often higher in conifer stands because of more leaf area indices. The differences are most pronounced during the dormant season when interception rates are low in hardwood stands. For example, two studies in the European forests have found that average yearly interception rates are around 25% for broadleaves species and about 45% for coniferous species (Augusto et al., 2002; Calder et al., 2003).

Canopy transpiration is often thought to increase asymptotically with leaf area index (L) for a species (Meinzer and Grantz, 1991; Raupach, 1995; Arneeth et al., 1996; Oren et al., 1996). It was assumed that annual transpiration does not differ considerably between broadleaf and coniferous forests (Roberts, 1983; Harding et al., 1992; Cannell, 1999). Large-scale afforestation resulted a rise in evapotranspiration, hence dropping in-stream flows (Farley et al., 2005; Sun et al., 2006, 2008), therefore impacting the effectiveness of water conservation strategy of plants at leaf or individual level. Quantifying the productivity-water loss tradeoffs at the ecosystem level is the primary parameter to analyze the carbon-water relationship in different forest types (Li et al., 2019). Many studies at ecosystem level have demonstrated that broadleaved forest have higher productivity and less water loss than needle-leave forests (Tan et al., 2015; Gower et al., 2001). This might be because deciduous leaves have higher rates of photosynthesis per unit leaf mass during favorable conditions than evergreens, given their higher leaf nitrogen content and specific leaf area, higher intrinsic photosynthetic capacity, and the reduced internal competition for light and carbon dioxide (Catovsky et al., 2002). For example, deciduous oaks compensate for having a shorter growing season by attaining a higher capacity to assimilate carbon for a given amount of intercepted solar radiation during the well-watered spring period. At saturating light levels, deciduous oaks gained carbon at six times the rate of evergreen oaks (Baldochi et al., 2010).

This water utilization behavior of the broadleaf and conifers may directly and significantly impact the hydrology and water yield of the forest. For example, the first study (Swank and Douglass, 1974) to examine differences in annual runoff and evapotranspiration (ET) between broadleaf and coniferous forests was performed in the United States using the paired-catchment method. Annual flow decreased with the conversion from broadleaf to coniferous forest. Changes in yearly runoff due to vegetation changes indicate changes in annual ET. Thus, the results indicate lower annual ET for broadleaf forests than for coniferous forests, suggesting that the presence of broadleaf forests is more beneficial from the viewpoint of water availability in downstream. In another study, long-term records of streamflow following the conversion of hardwood stands to conifers show reduced water yield (Hornbeck et al., 1997; Komatsu et al., 2009). For example, this trend has been instigated in Japan (Komatsu et al., 2009). More extensive evidence of the lower annual evapotranspiration for broadleaf forests compared to coniferous forests was provided by surveying the results of numerous paired-catchment studies (Bosch and Hewlett, 1982). A very latest survey of a global dataset also demonstrated that ET of broadleaf forests is lower than coniferous, resulting in a higher annual runoff for broadleaf. The study also suggested that this condition is only valid for broadleaf deciduous forests (Komatsu et al., 2011). The differences between conifers and deciduous trees are often incorporated into large-scale models because of differences in xylem anatomy (vessels versus tracheids), leaf

longevity, leaf area index, and growing season (Roberts and Rosier, 2005). Therefore, from all these survey results it is concluding that the increase in annual runoff due to deforestation tended to be lower of broadleaf forests than coniferous forests, which suggests the generality of the yearly ET for gaining growth and productivity of broadleaf forests is lower than coniferous forests.

Conclusion

In the past, many afforestation projects were established without knowing their carbon and water interaction. The difference of change in water yield (%), as well as an increase in streamflow (mm) after harvesting of broadleaved and needle-leaved forest in low and high precipitation regions, along with other adjoining factors are giving us a clue as to how the future afforestation policy needs to be revised. When, where, and why to plant/cut the tree is important questions to address. The results showed that needle-leaved forests in lower precipitation regions are expected to consume more water than broadleaved in higher rainfall regions. This study can speculate from these results that afforestation and conversion of broadleaved to conifers or mixed in higher precipitation regions might be more useful to get maximum productivity. Conversely, in lower precipitation regions scattered plantation of broadleaved primarily deciduous species along with shrubs and grasses might be an option to maintain the carbon and water tradeoff of global forests. However, sustainable forest management and targeted planning for establishment of future plantations need to take into account a broader prospective of multifunctional objectives is prerequisite to mitigate the future challenges of climate change.

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Conflict of Interests. Authors declare that there is no conflict of interests.

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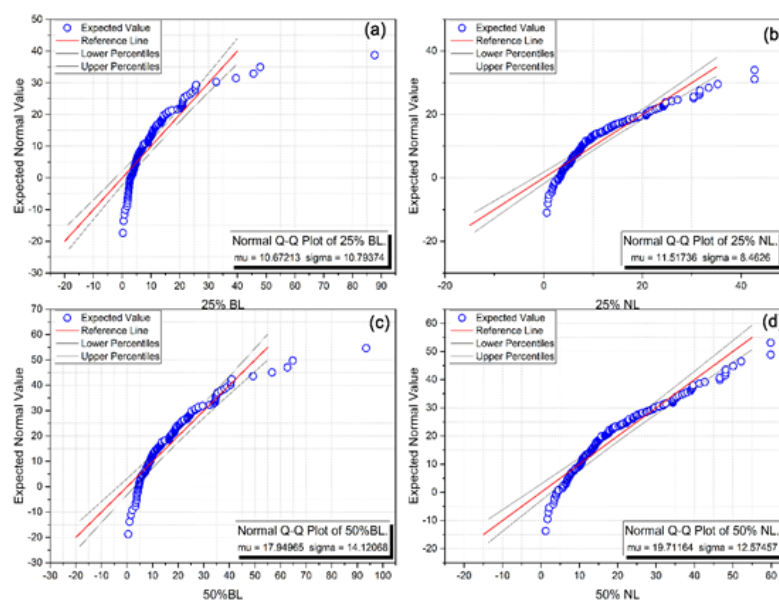
APPENDIX

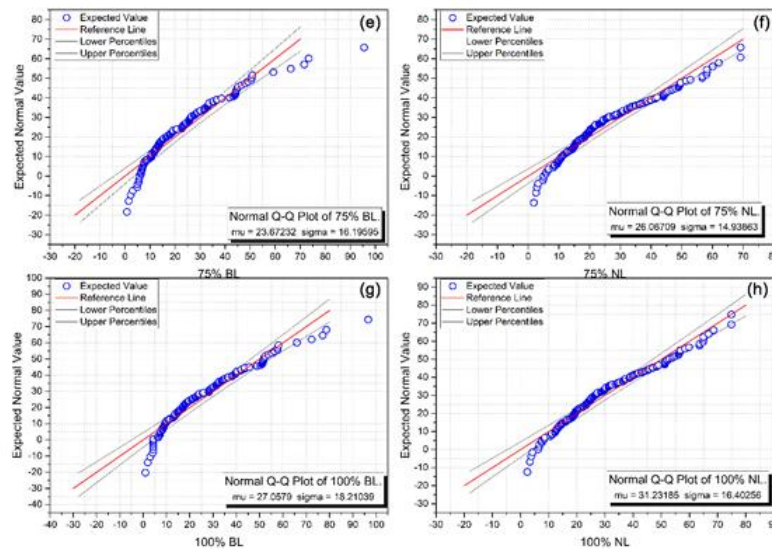
Appendix 1. Normality test results of Shapiro-wilk reject normality at $P < 0.05$

H.I.	DF	Statistics	P-value	Decision at level (5%)
25%NL	167	0.87663	1.6514E-10	Reject normality
25%BL	133	0.6918	2.38698E-15	Reject normality
50%NL	167	0.92063	6.51603E-8	Reject normality
50%BL	133	0.83991	1.02842E-10	Reject normality
75%NL	167	0.9465	5.95852E-6	Reject normality
75%BL	133	0.90299	8.54516E-8	Reject normality
100%NL	167	0.96282	1.93127E-4	Reject normality
100%BL	133	0.92935	3.17434E-6	Reject normality

H.I.: harvesting intensity (%)

Appendix 2. QQ plot representing the distribution of change in water yield (%) after 25, 50, 75 and 100% of harvesting intensities of the dataset. The scores are negatively skewed (fewer scores at the low end)





Appendix 3. Kruskal-Wallis ANOVA, representing the descriptive statistic results of different harvesting intensities of broadleaved (BL) and needle leaved (NL) forests on change in water yield (%)

H.I.	N	Min	Q1	Median	Q3	Max
“25%NL”	167	0.61635	5.71968	9.04704	15.52163	42.71476
“25%BL”	133	0.25043	4.24867	7.88758	13.50434	87.69458
“50%NL”	167	1.22515	10.82046	16.59292	26.87225	59.86033
“50%BL”	133	0.4996	8.15101	14.62185	23.79504	93.44391
“75%NL”	167	1.82653	15.39764	22.98264	35.53398	69.10673
“75%BL”	133	0.74753	11.74769	20.43853	31.89731	95.53162
“100%NL”	167	2.42063	19.5279	28.463	42.36111	42.36111
“100%BL”	133	0.99423	13.54604	22.96223	38.15074	96.61086

H.I: harvesting intensity (%), $P < 0.0001$, Chi-square = 295.027

Appendix 4. Wilcoxon signed ranks test

Paired sample	W	Z	P-value
“25%BL”-“25%NL”	5869	3.17332	0.00151
“50%BL”-“50%NL”	5903	3.24968	0.00116
“75%BL”-“75%NL”	5927	3.30358	9.54597E-4
“100%BL”-“100%NL”	6206	3.93016	8.48907E-5

* $P < 0.05$

Appendix 5. Data set used in the synthesis

Broadleaved forest								
Source	Catchment	Country	Elv. (m)	Soil type	Area (ha)	MAP (mm)	MAS (mm)	DF (%)
Bosch and Hewlett (1982)	Coweeta 13	USA	810	Sandy clay loam	16	1900	889	100

Bosch and Hewlett (1982)	Coweeta 19	USA	960	Sandy clay loam	28	2001	1222	22
Bosch and Hewlett (1982)	Coweeta 1	USA	840	Sandy clay loam	16	1725	739	100
Bosch and Hewlett (1982)	Coweeta 28	USA	1200	Sandy clay loam	144	2270	1532	65
Bosch and Hewlett (1982)	Coweeta 17	USA	885	Sandy clay loam	14	1895	775	100
Bosch and Hewlett (1982)	Coweeta 22	USA	1035	Sandy clay loam	34	2068	1275	50
Bosch and Hewlett (1982)	Coweeta 3	USA	825	Sandy clay loam	9	1814	607	100
Bosch and Hewlett (1982)	Coweeta 10	USA	975	Sandy clay loam	86	1854	1072	30
Bosch and Hewlett (1982)	Coweeta 41	USA	1065	Sandy clay loam	29	2029	1285	53
Bosch and Hewlett (1982)	Coweeta 6	USA	793	Sandy clay loam	9	1854	838	80
Bosch and Hewlett (1982)	Kericho Sambret	Kenya	2200	Deep friable clay	688	1905	416	34
Bosch and Hewlett (1982)	Kimakia A	Kenya	2440	Deep friable clay	35	2014	568	100
Bosch and Hewlett (1982)	Fernow 1	USA	755	Stony silt loam	30	1524	584	85
Bosch and Hewlett (1982)	Fernow 2	USA	780	Stony silt loam	15	1500	660	36
Bosch and Hewlett (1982)	Fernow 5	USA	780	Stony silt loam	36	1473	732	20
Bosch and Hewlett (1982)	Fernow 3	USA	805	Stony silt loam	34	1500	607	13
Bosch and Hewlett (1982)	Fernow 7	USA	800	Stony silt loam	24	1469	788	50
Bosch and Hewlett (1982)	Fernow 6	USA		Stony silt loam	22	1440	493	50
Bosch and Hewlett (1982)	Leading Ridge WS2	USA	385	Silt loam	43	1004	321	20
Bosch and Hewlett (1982)	Placer County Ws C	USA	168	Clay loam	5	635	145	99
Bosch and Hewlett (1982)	Maimai M7	New Zealand	300	Stoney silt loam	4	2600	1500	100
Bosch and Hewlett (1982)	Maimai M9	New Zealand	310	Stoney silt loam	8	2600	1500	75
Andréassian (2004)	Leading Ridge 2	USA			43	1060	440	86
Andréassian (2004)	Dantzoud	Armenia			14100	680	413	11
Andréassian (2004)	Girants	Armenia			12200	700	224	7
Bent (2001)	Dickey brook	USA	308		308	1250	430	32
Brechtel and Fuhrer(1991)	Krofdorf A1	Germany	336	Rocky	9.3	650	300	100
Fahey and Jackson (1997)	Big bush DC1	New Zealand			8.6	1530	610	83
Fahey and Jackson (1997)	Big bush DC4	New Zealand			20.2	1530	670	94
Fritsch (1992)	Hakhoum	Armenia			####	675	268	7
Sahin and Hall (1996)	WS2L.R.	USA	360	Silt loam	43	1060	440	43
Sahin and Hall (1996)	WS4H.B	USA	606	Sandy loam	36	1340	860	33
Stednick (1996)	Coweeta 7	USA	900	Loam	59	1825	1140	100
Stednick (1996)	Fernow 3	USA	805	Silt loam	34	1500	610	91
Stednick (1996)	Ouachita, OKWS10	USA		Loam	5.7	1317	1652	50
Stednick (1996)	Ouachita, OKWS12	USA		Loam	5.9	1317	1652	100
Stednick (1996)	Ouachita, OKWS14	USA		Loam	4.3	1317	1652	50

Stednick (1996)	Ouachita, OKWS15	USA		Loam	5.1	1317	1652	100
Stednick (1996)	Ouachita, OKWS17	USA		Loam	4.2	1317	1652	50
Stednick (1996)	Ouachita, OKWS18	USA			4.1	1317	1652	100
Anderson et al. (1976)	WN-Carolina 1	USA			15	1828	787	100
Anderson et al. (1976)	WN-Carolina 3	USA			9	1803	610	100
Anderson et al. (1976)	WN-Carolina 5	USA			28	2006	1219	22
Anderson et al. (1976)	WN-Carolina 6	USA			83	1854	1067	30
Anderson et al. (1976)	WN-Carolina 7	USA			28	2032	1295	35
Anderson et al. (1976)	NW-Virginia 1	USA			22	1448	762	100
Anderson et al. (1976)	NW-Virginia 3	USA			23	1447	762	50
Anderson et al. (1976)	NW-Virginia 5	USA			33.4	1498	635	14
Hornbeck et al. (1993), Kabeya et al. (2015)	Pennsylvania LR-WS3	USA	340		104	1060	440	43
Hornbeck et al. (1993), Kabeya et al. (2015)	Pennsylvania LR-WS2	USA	360		43	1060	440	24
Swift and Swank (1981)	Coweeta 13	USA	810	Clay loam	16	1900	889	100
Swift and Swank (1981)	Coweeta 37	USA	1300	Sandy clay loam	44	2220	1604	100
Swift and Swank (1981)	Coweeta 28	USA	1200		144	2320	1534	65
Andréassian (2004)	Karuah/Kokata	Australia			97.4	1565	531	29
Andréassian (2004)	Karuah/Coachwood	Australia			37.5	1447	362	61
Andréassian (2004)	Karuah/Corkwood	Australia			41.1	1636	505	40
Andréassian (2004)	Karuah/Jackwood	Australia			12.5	1373	311	79
Andréassian (2004)	Karuah/Bollygum	Australia			15.1	1518	505	32
Andréassian (2004)	Monda 1	Australia		Rocky Krasnozems	6.3	1876	702	75
Andréassian (2004)	Monda 2	Australia		Rocky Krasnozems	4	1813	550	75
Andréassian (2004)	Monda 3	Australia		Rocky Krasnozems	7.3	1763	632	75
Andréassian (2004)	Myrtle 2	Australia		Rocky Krasnozems	30.5	1590	852	74
Andréassian (2004)	Picaninny	Australia			53	1180	332	78
Andréassian (2004)	Black Spur 1	Australia		Rocky Krasnozems	17	1652	504	60
Andréassian (2004)	Black Spur 3	Australia		Rocky Krasnozems	7.7	1612	530	60
Andréassian (2004)	Wicksend	Australia			68	1200	440	22
Andréassian (2004)	Wilbob	Australia			86	1200	392	12
Andréassian (2004)	Clem creek	Australia		Rocky clay loam	46.4	1445	190	95
Andréassian (2004)	Yarragil 4L	Australia			126	1120	4.3	66
Sahin and Hall (1996)	Hansen	Australia		Gravel	80	1200	232	75
Komatsu et al. (2011), Pearce et al. (1980)	Maimai M7	New Zealand				2600	1550	100
Pearce et al. (1980)	Maimai M9	New Zealand				2600	1550	75
Stednick (1996)	Fernow 3	USA	805	Silt loam	34	1500	610	13
Stednick (1996)	Fernow 5	USA	760	Silt loam	36	1470	760	20
Stednick (1996)	Fernow 6	USA		Silt loam	22	1440	490	50
Stednick (1996)	Fernow 7	USA	800	Silt loam	24	1470	790	50
Stednick (1996)	Leading Ridge PA2	USA	358	Silt loam	43	1000	320	20
Stednick (1996)	Coweeta, NC7	USA	900	Loam	59	1825	1140	100
Stednick (1996)	Grant forest GA18	USA	165	Sandy loam	33	1220	470	100
Stednick (1996)	Ouachita, OKWS10	USA		Loam	6	1317	1652	50
Stednick (1996)	Ouachita, OKWS12	USA		Loam	6	1317	1652	100
Stednick (1996)	Ouachita, OKWS14	USA		Loam	4	1317	1652	50
Stednick (1996)	Ouachita, OKWS15	USA		Loam	4	1317	1652	100
Stednick (1996)	Ouachita, OKWS17	USA		Loam	4	1317	1652	50

Needle-leaved forest								
Source	Catchment	Country	Elev. (m)	Soil type	Area (ha)	MAP (mm)	MAS (mm)	DF (%)
Bosch and Hewlett (1982)	Needle Branch	USA	312	Sand stone	71	2483	1885	82
Bosch and Hewlett (1982)	Deer Creek	USA	312	Sand stone	303	2474	1906	25
Bosch and Hewlett (1982)	H.J. Andrews 1	USA	700	Clay loams	96	2388	1376	100
Bosch and Hewlett (1982)	H.J. Andrews 3	USA	760	Clay loams	101	2388	1346	30
Bosch and Hewlett (1982)	H.J. Andrews 6	USA	900	Volcaniclastics	13	2150	1290	100
Bosch and Hewlett (1982)	H.J. Andrews 7	USA	900	Volcaniclastics	21	2150	1290	60
Bosch and Hewlett (1982)	H.J. Andrews 10	USA	500	Volcaniclastics	9	2330	1650	100
Bosch and Hewlett (1982)	Coyote Creek 1	USA	901	Gravelly loam	59	1230	627	50
Bosch and Hewlett (1982)	Coyote Creek 2	USA	901	Gravelly loam	68	1230	630	30
Bosch and Hewlett (1982)	Coyote Creek 3	USA		Gravelly loam	50	1230	630	100
Bosch and Hewlett (1982)	Workman Creek, NF	USA	2225	Clay loam	100	813	86	73
Bosch and Hewlett (1982)	Workman Creek, SF	USA	2165	Clay loam	129	813	87	83
Bosch and Hewlett (1982)	Fool Creek	USA	3200	Permeable soil	289	762	283	40
Bosch and Hewlett (1982)	Castle Creek	USA	8207	Soil of igneous origin	364	639	71	17
Bosch and Hewlett (1982)	Beaver Creek 1	USA	1700	Stony clay	124	457	24	100
Bosch and Hewlett (1982)	Beaver Creek 3	USA	1600	Stony clay	146	457	18	83
Bosch and Hewlett (1982)	Wagon Wheel Gap	USA	3110	Rocky clay loam	81	536	157	100
Burton (1997)	Brownie Creek	USA	3082	Sand stone	2134	787	300	25
Troendle et al. (2001), Pike and Scherer (2003)	Coon creek	USA			1673	870	440	24
Cosandey (1990)	Latte	France			20	1900	1278	100
Stednick (1996)	Workman Ce.AZ	USA	2225	Clay loam	100	833	86	32
Stednick (1996)	N.Fork	USA	2225	Clay loam	100	810	86	32
Stednick (1996)	Wagonwheel Gap.CO	USA	3110	Rocky clay loam	81	544	157	100
Stednick (1996)	Chicken Creek M.OR1	USA	1523	Ash		1355	472	50
Stednick (1996)	Chicken Creek M.OR2	USA	1523	Ash		1355	460	50
Stednick (1996)	Chicken Creek M.OR3	USA	1523	Ash		1355	372	50
Stednick (1996)	Fool Creek, CO	USA	3200	Granite	289	760	280	40
Stednick (1996)	Fraser Forest, CO	USA	3200	Granite	289	712	283	66
Stednick (1996)	Deadhorse Cr. CO	USA	3120	Granite	270	762	500	36
Stednick (1996)	White Spar C	USA	1420	Quartz	5	450	43	100
Stednick (1996)	Castle Creek, AZ	USA		Igneous	364	640	71	17
Stednick (1996)	Deer Creek, OR	USA	312	Marine sand stone	303	2480	1910	25
Stednick (1996)	Needle Branch, OR	USA	312	Perm sand stone	71	2480	1885	82
Stednick (1996)	Blue Mts1	USA	1523	Ash		1355	472	50
Stednick (1996)	Blue Mts2	USA	1523	Ash		1355	460	50
Stednick (1996)	Blue Mts3	USA	1523	Ash		1355	372	50
Stednick (1996)	St Louis creek	USA	3200	Granite	289	712	283	100

Stednick (1996)	Thomas creek, AZ	USA	2600	Loamy	227	768	500	34
Stednick (1996)	Willow creek, AZ	USA		Loam		749	512	62
Anderson (1976)	Western Oregon 1	USA			93	2362	1447	100
Anderson (1976)	Colorado 2	USA			281	762	279	40
Anderson (1976)	Arizona 2	USA			100	812	86	32
Cheng (1989)	Camp Creek, BC	USA	1920	Granite	3390	600	140	30
Cheng (1989)	Hinton, Alberta	Canada			1497	513	147	50
Cheng (1989)	Cabin Creek, Alberta	Canada			212	840	310	21
Scott et al. (2000)	Biesieviei	South Africa	580		27.2	1298	593.6	100
Scott et al. (2000)	Bosboukloof	South Africa	671		200.9	1564	245.9	100
Scott et al. (2000)	Witklip-6	South Africa	1080		165.3	929	259.7	100
Scott et al. (2000)	WitkliP-5	South Africa	1340		108	929	261.7	51
Webb (2009)	Canobolas A	Australia	1200		55.3	1080	289	100
Webb (2009)	Canobolas B	Australia	1180		55.4	1080	247	100
Cosandey (1990)	Latte	France			20	1900	1278	100
Adams and Flower (2006)	Maimai M5	New Zealand		Gritty silt loam	2.31	2453	1578	100
Adams and Flower (2006)	Maimai M8	New Zealand		Gritty silt loam	3.84	2453	1213	95
Adams and Flower (2006), Rowe et al. (2002)	Glenbervie, Logbridge	New Zealand			12.6	1920	830	100
Adams and Flower (2006), Rowe et al. (2002)	Glenbervie, Pines	New Zealand			15.5	1920	760	100
Adams and Flower (2006), Rowe et al. (2002)	Moumoukai, Central	New Zealand		Clay loam	11.42	1690	660	100
Komatsu et al. (2011), Adams and Flower (2006)	Moumoukai, South	New Zealand		Clay loam	14.98	1700	646	100
Adams and Flower (2006), Rowe et al. (2002)	Purukohukohu, Puruki	New Zealand		Sandy loam to loamy sand	34.4	1590	540	100
Adams and Flower (2006), Rowe et al. (2002)	Pakuratahi	New Zealand		Silt loam	345	1097	380	87
Adams and Flower (2006)	Moutere, C13	New Zealand			7.65	1010	64	100

MAP: mean annual precipitation, MAS: mean annual streamflow, DF: deforestation, Elv.: elevation